



(Strike-slip)

Faults & shear zones: strain localization in nature

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Géosciences Montpellier



Faults: highly localized deformation @ the surface



*2001/11/14 Kokoxili earthquake
Kunlun fault, photo R. Lacassin*



San Andreas, California, photo J.C. Bousquet

Geodetic observations after major earthquakes:

ductile flow beneath the fault

Pollitz et al. (2000, 2001, 2003), Montési (2004, 2005)...

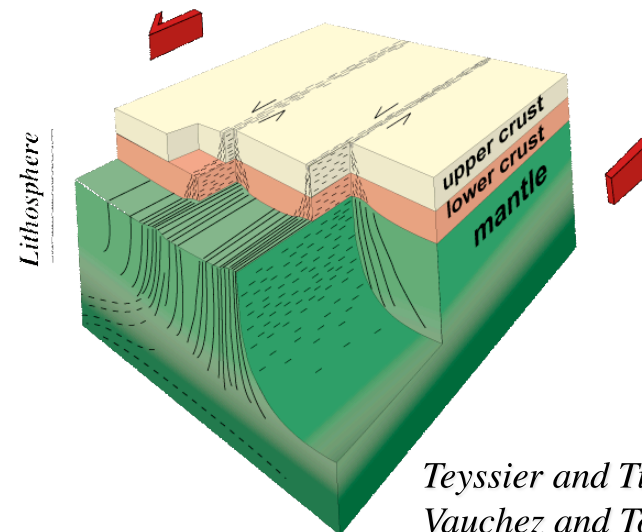
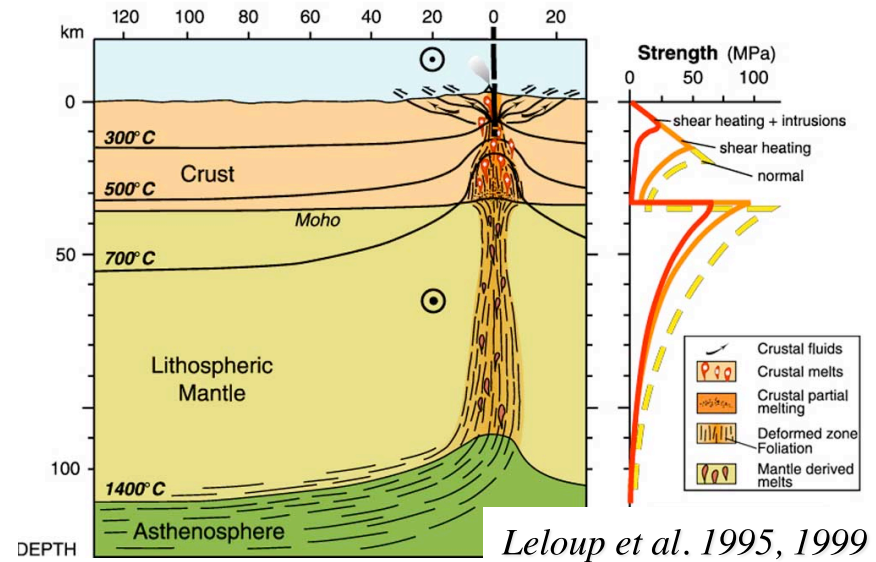
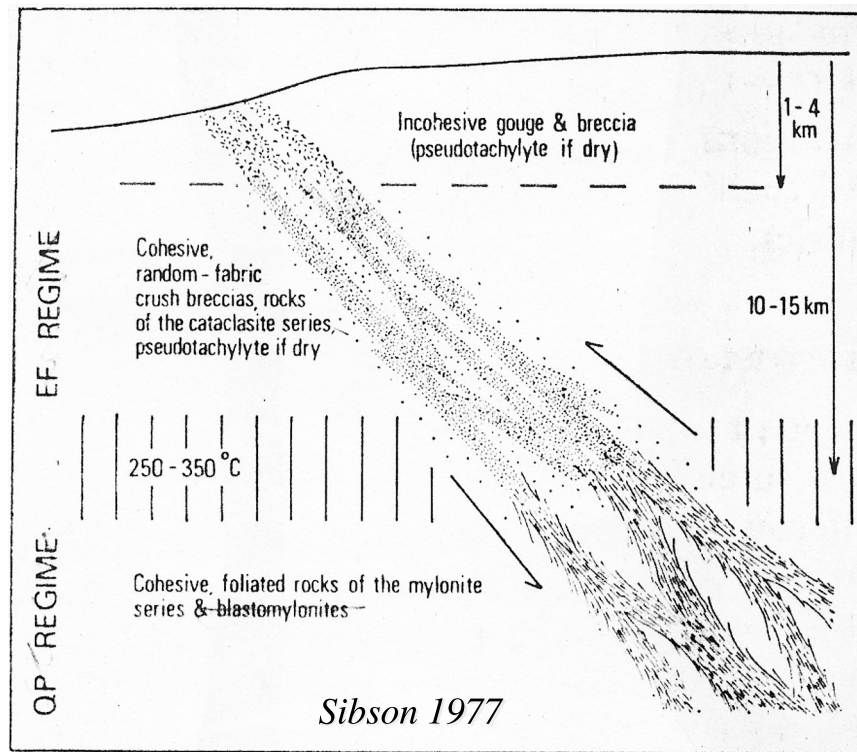
***Mantle Flow Beneath a Continental
Strike-Slip Fault: Postseismic
Deformation After the 1999
Hector Mine Earthquake***

Fred F. Pollitz, Chuck Wicks, Wayne Thatcher
SCIENCE VOL 293 7 SEPTEMBER 2001*

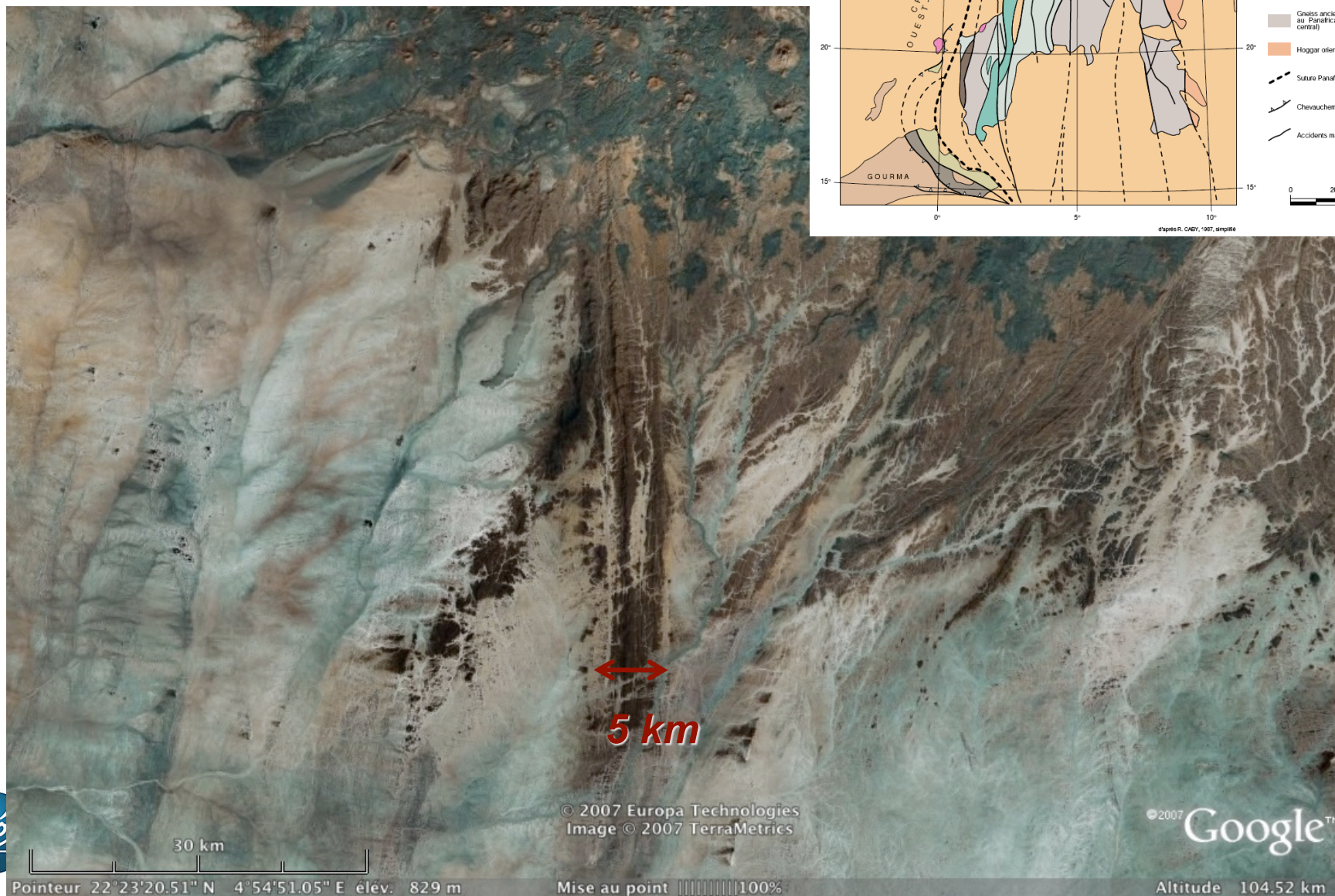
***How deep do these faults penetrate within the Earth?
How does the deformation (strain distribution) vary with depth ?***



How deep do these faults penetrate within the Earth? How does the deformation (strain distribution) vary with depth ?



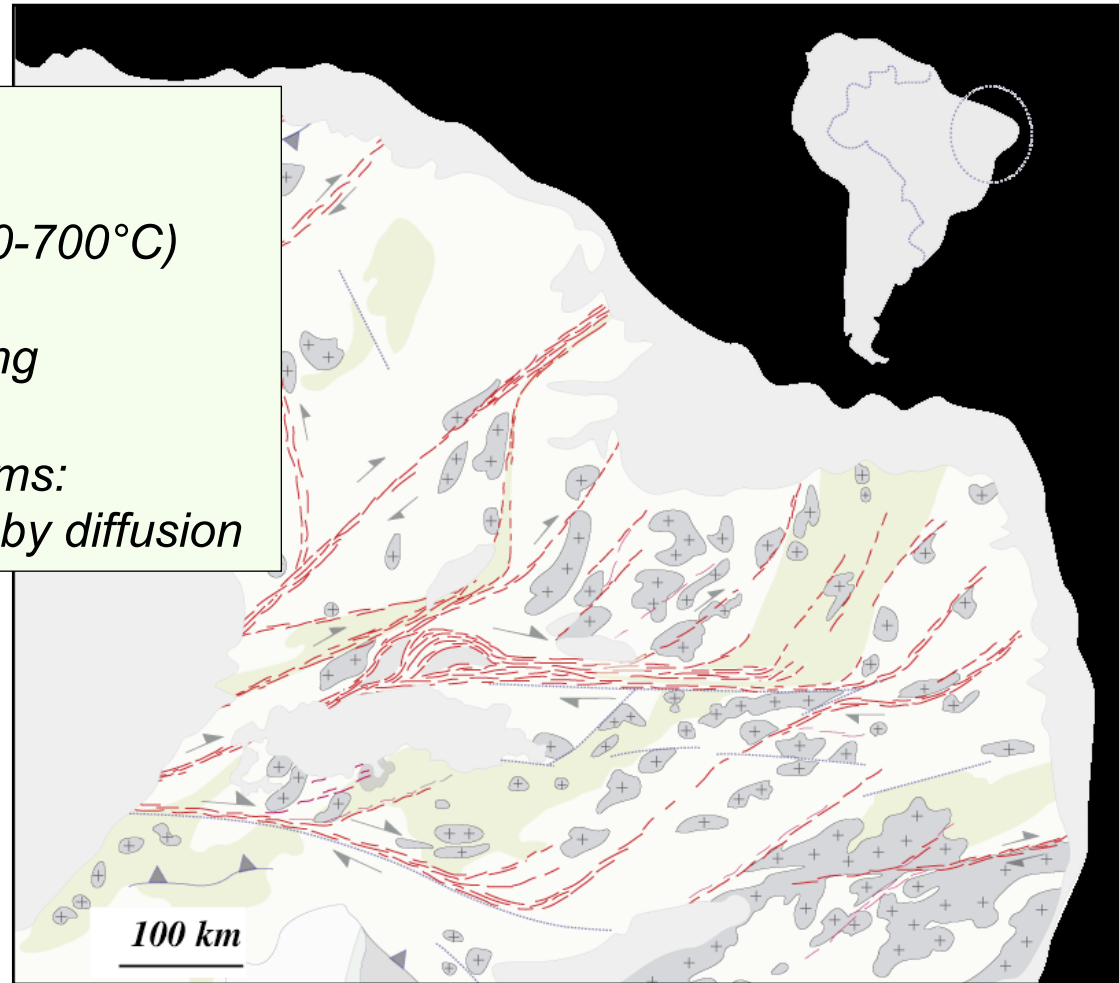
In the upper to middle crust: the 4°50 shear zone in Hoggar



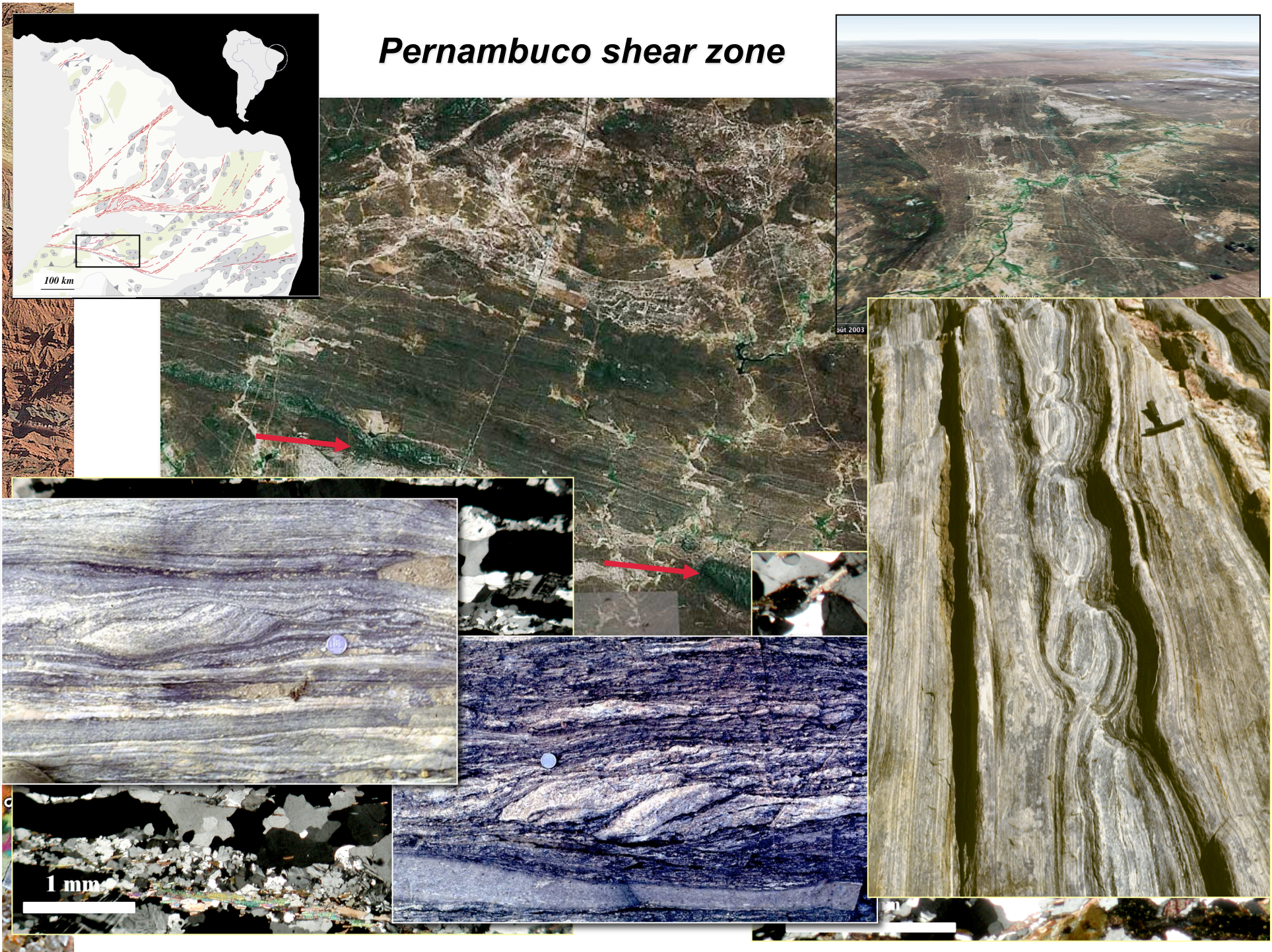
Wrench faults in the middle to lower crust the Borborema shear zone system (NE Brazil): localized strike-slip deformation in a partially molten crust

Northeast Brazil:

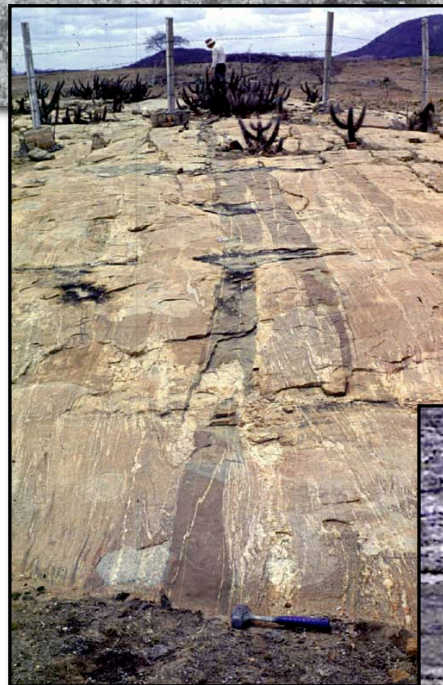
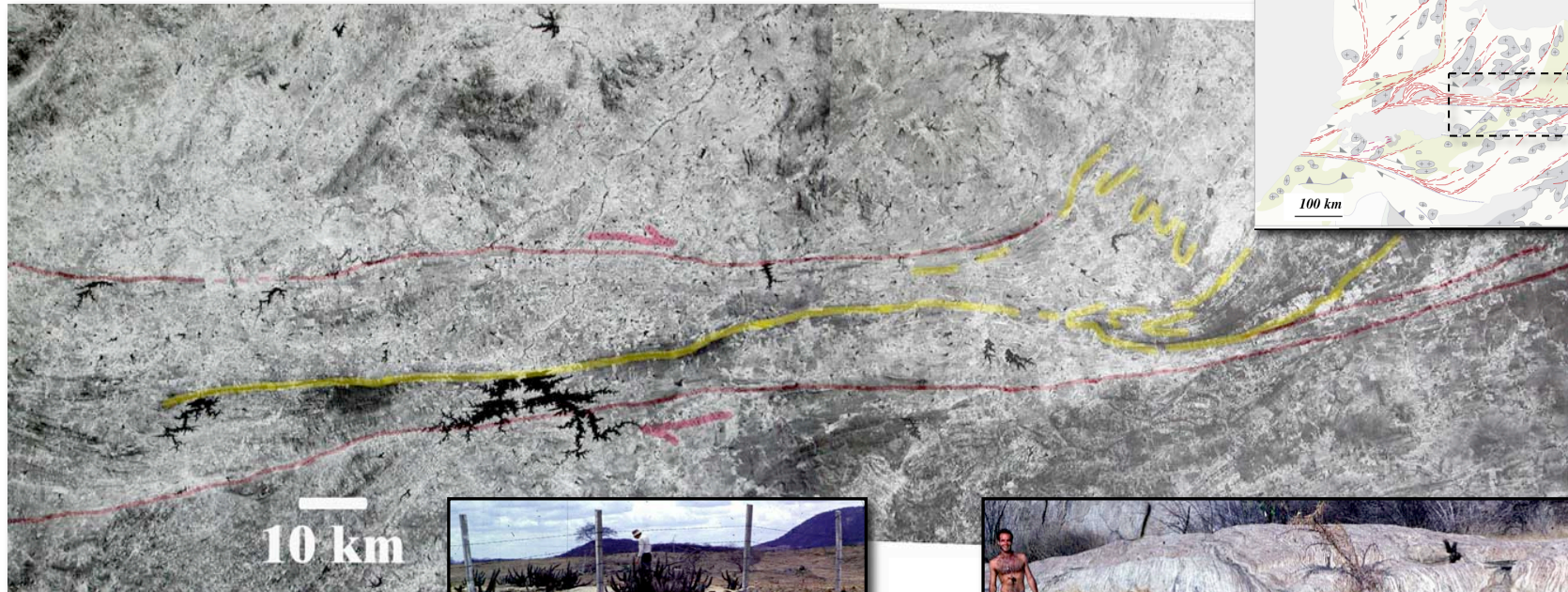
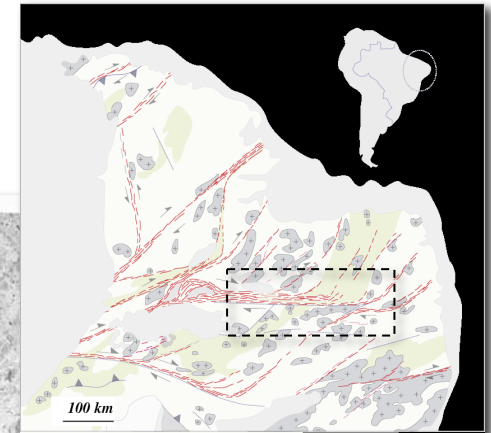
- Middle crust (15-17km)
- HT-LP metamorphism (650-700°C)
- Complex geometry
- Synkinematic partial melting and magmatism
- HT deformation mechanisms:
dislocation creep assisted by diffusion



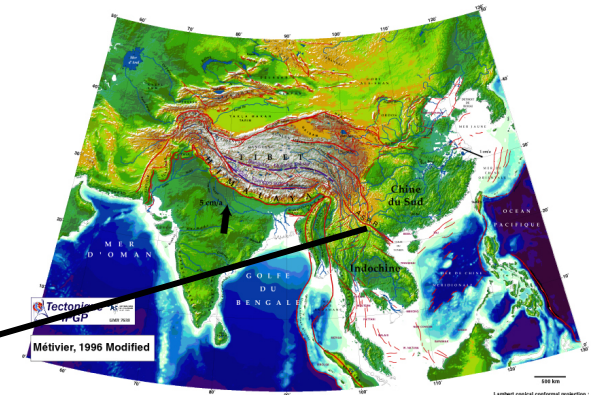
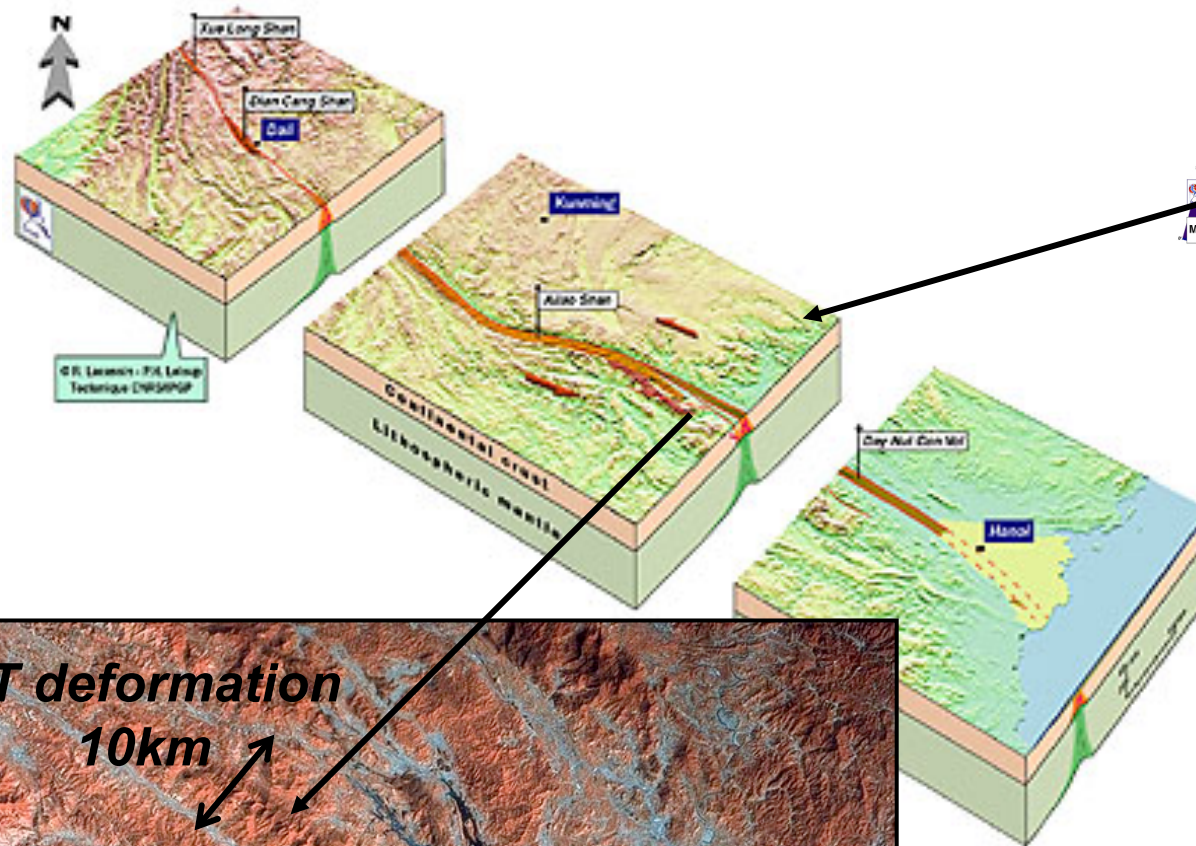
Pernambuco shear zone



Borborema shear zone system, NE Brazil

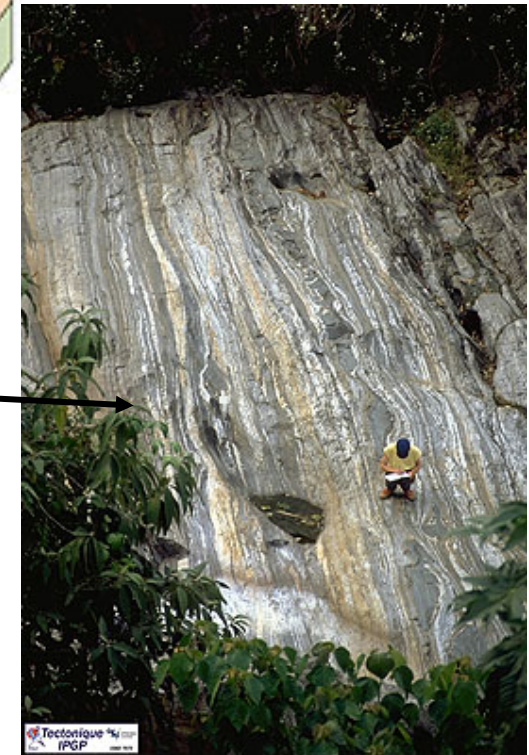


the Ailao Shan - Red River shear zone



HT deformation
10km

LT deformation
<1 km



Middle to lower crust
Além-Paraíba shear zone
(SE Brazil):

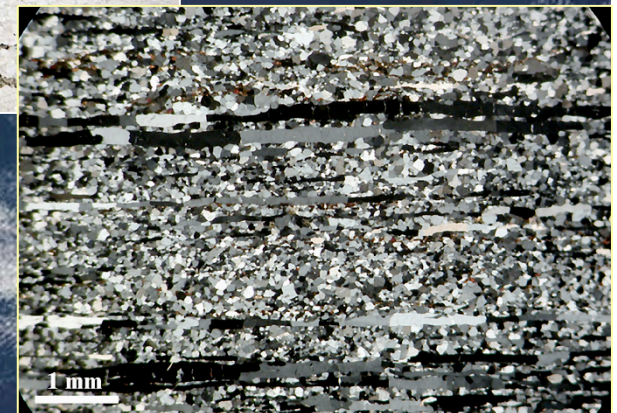
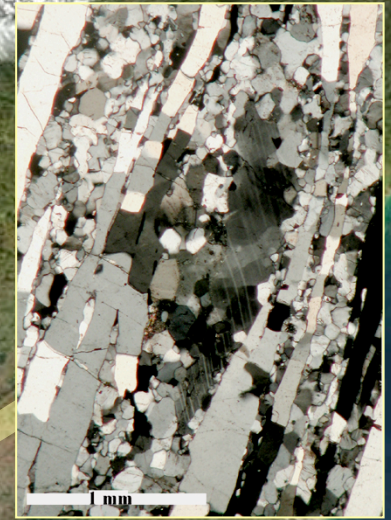
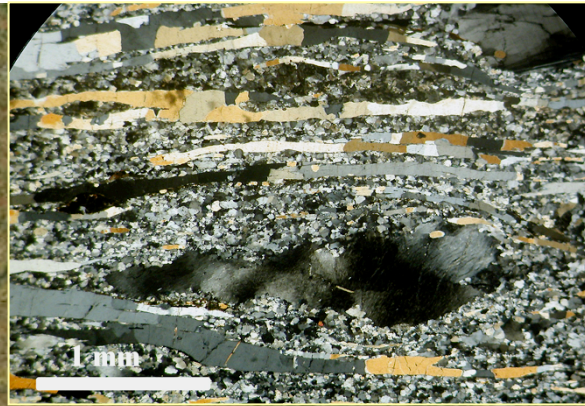
- 20-25km deep
- <5km wide
- HT-LP (750°C)
- Mostly mafic granulite
- No in-situ partial melting
- Synkinematic magmatism
- HT deformation mechanisms:
dislocation creep + diffusion



***Middle to lower crust
Além-Paraíba shear zone
(SE Brazil):***

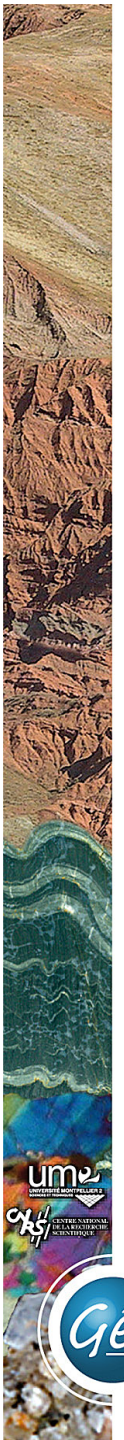
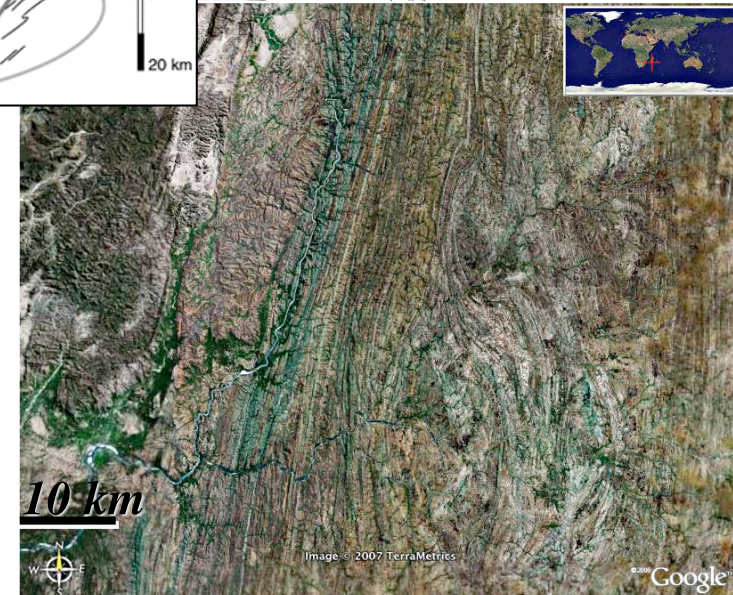
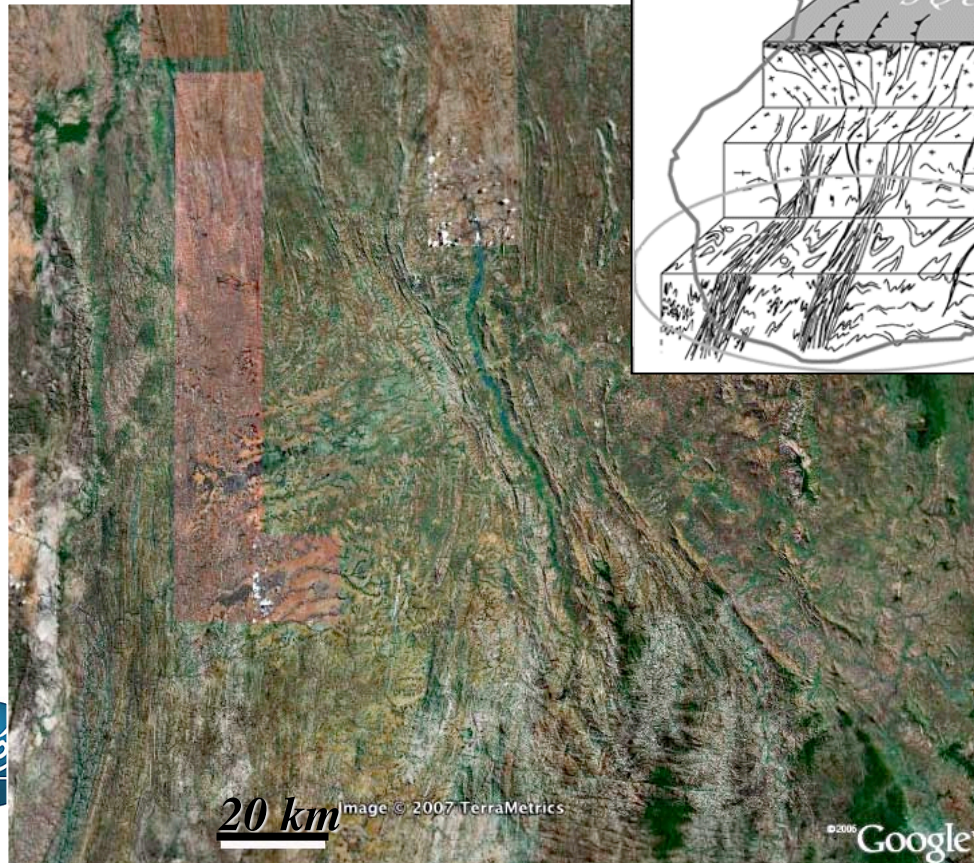
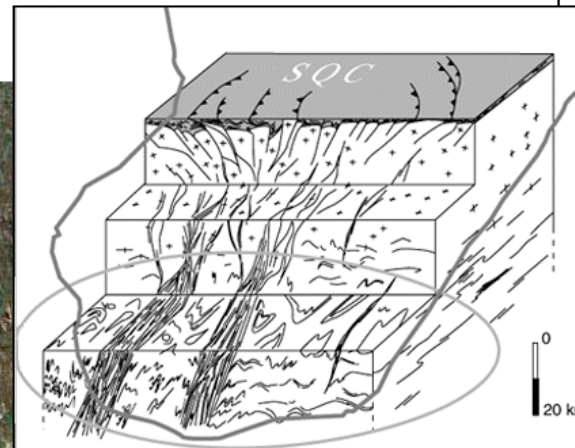
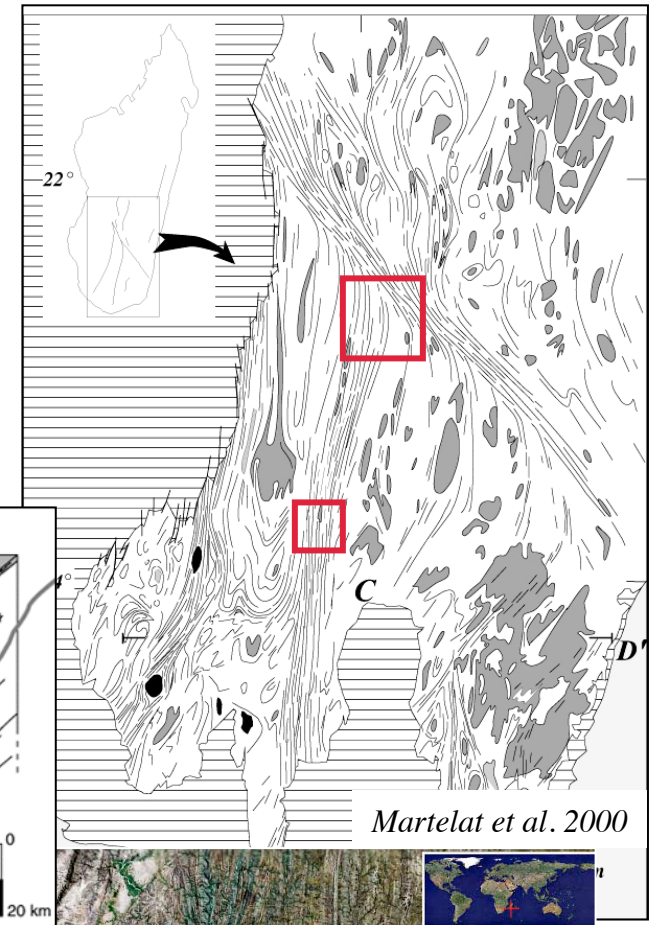
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*mechanisms:
dislocation creep +
diffusion*



Madagascar shear zone system:

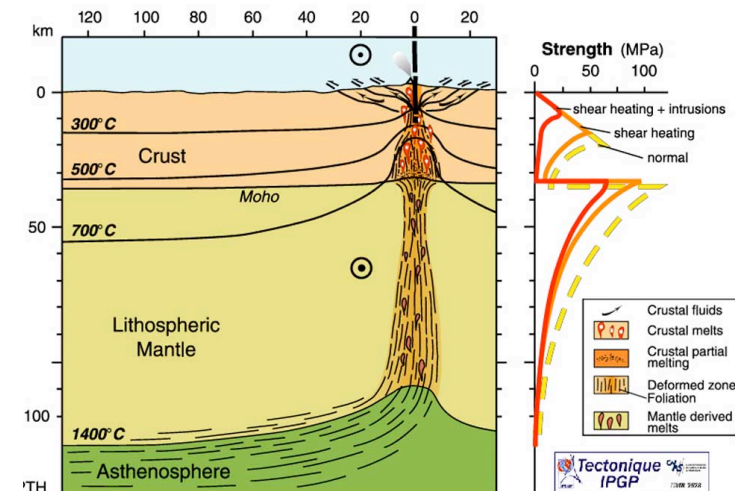
- Felsic middle crust to mafic lower crust
- Faults up to 30km wide
- HT ($750^{\circ}\text{C} \pm 50^{\circ}\text{C}$), low to interm P
- Synkinematic melting and magmatism
- HT deformation mechanisms:
dislocation creep + diffusion



Large-scale wrench faults in the ductile crust

Strain localization even at partial melting temperatures, but not very efficient:

- up to 30km wide shear zones
- complex geometries



- Deformation under HT, low stress conditions:
- “Coarse-grained” mylonites : dislocation creep, but fast diffusion (grain growth)
- "intrinsic" strain localization processes (shear heating or grain size reduction) not effective, with exception of mechanical anisotropy due to CPO or compositional layering

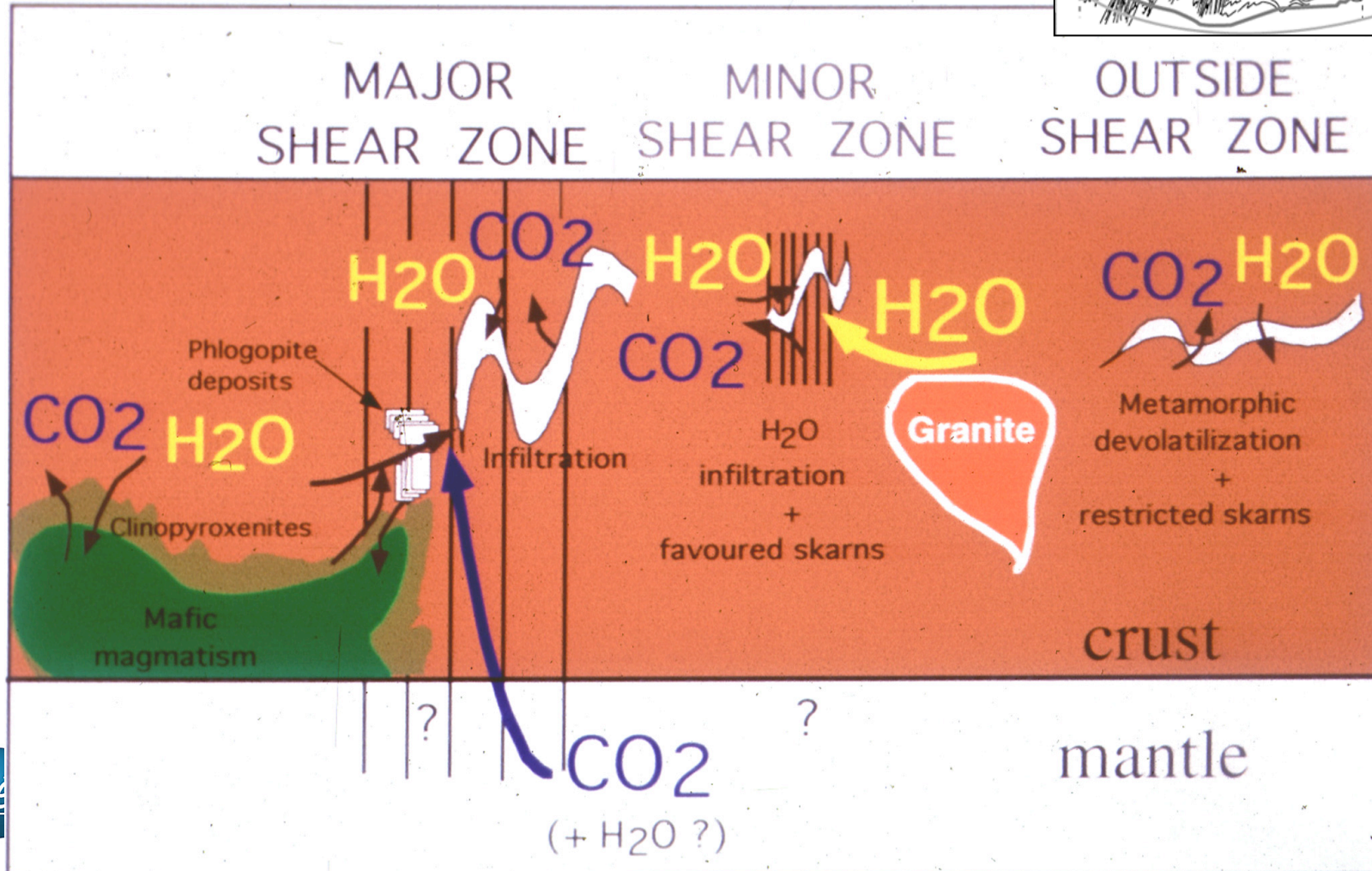
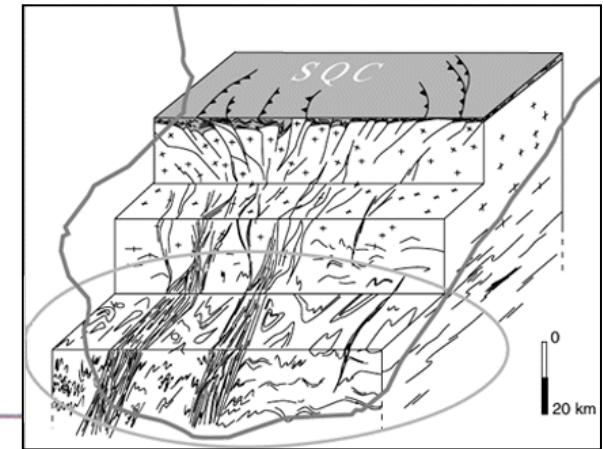
Which processes produce strain localization?

- coupling to shallow localized deformation
- channeling of fluids (magmas, CO₂..)
- rheological heterogeneity (lateral gradients in *T* or in composition)

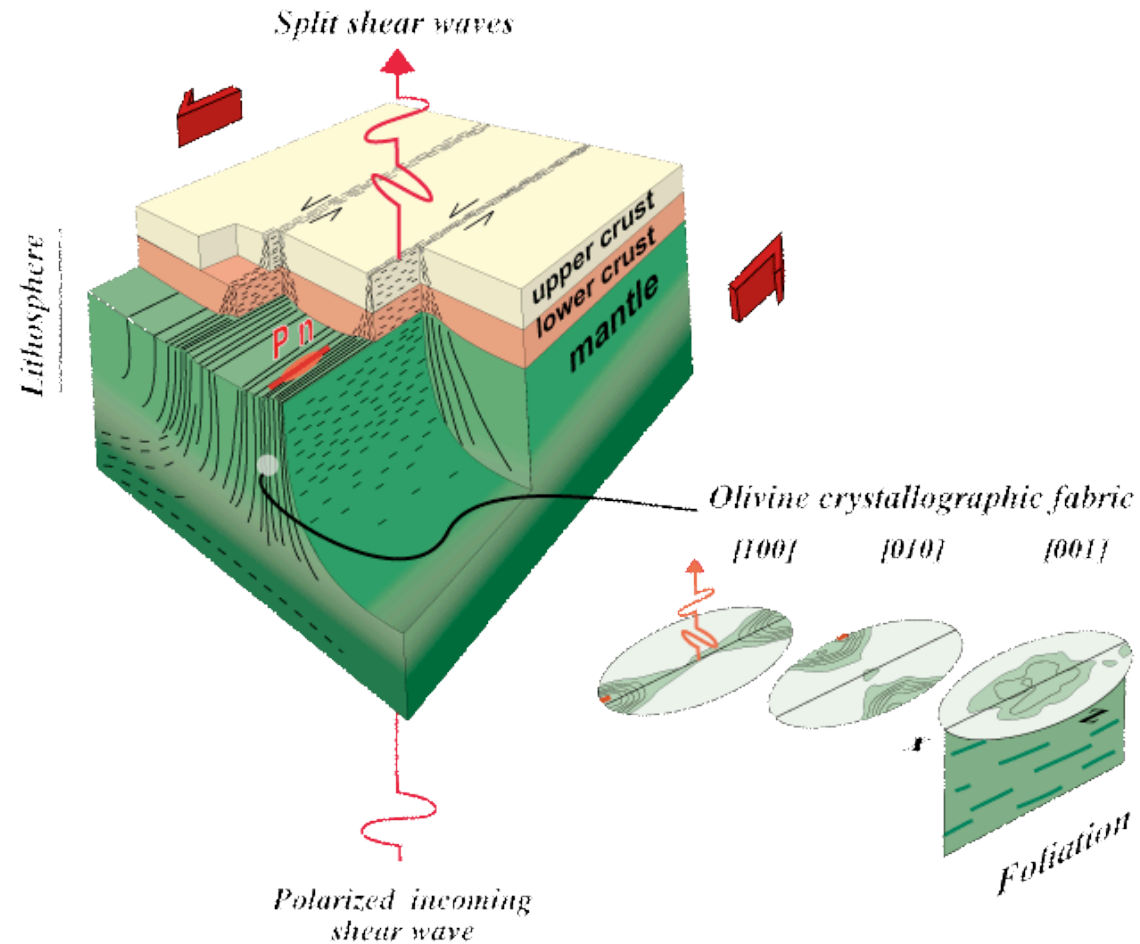


Do these large-scale faults penetrate in the mantle?

Madagascar shear zone system



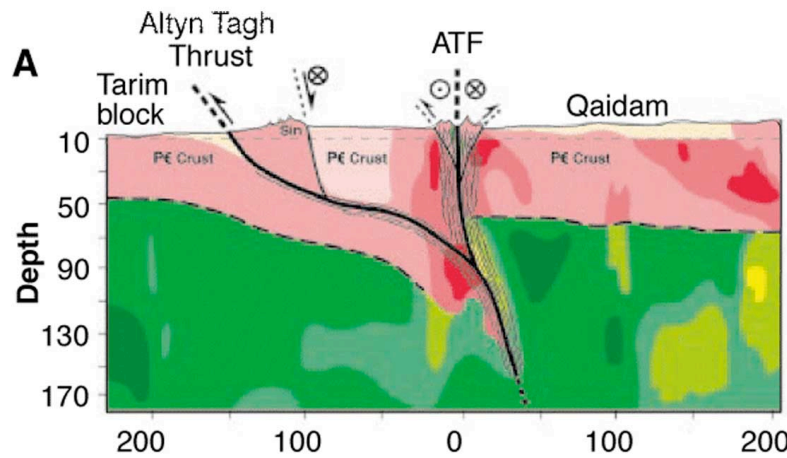
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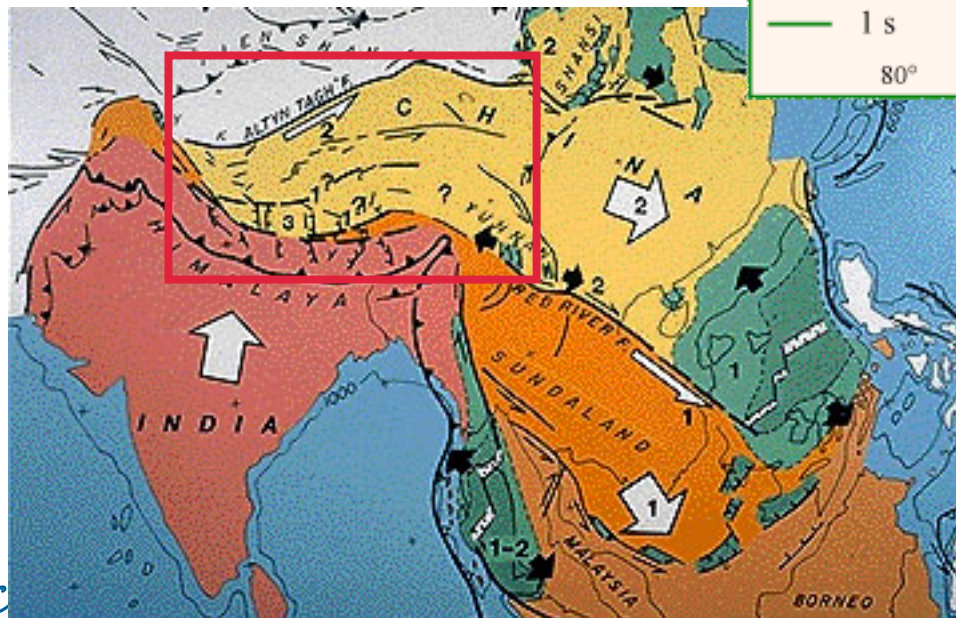
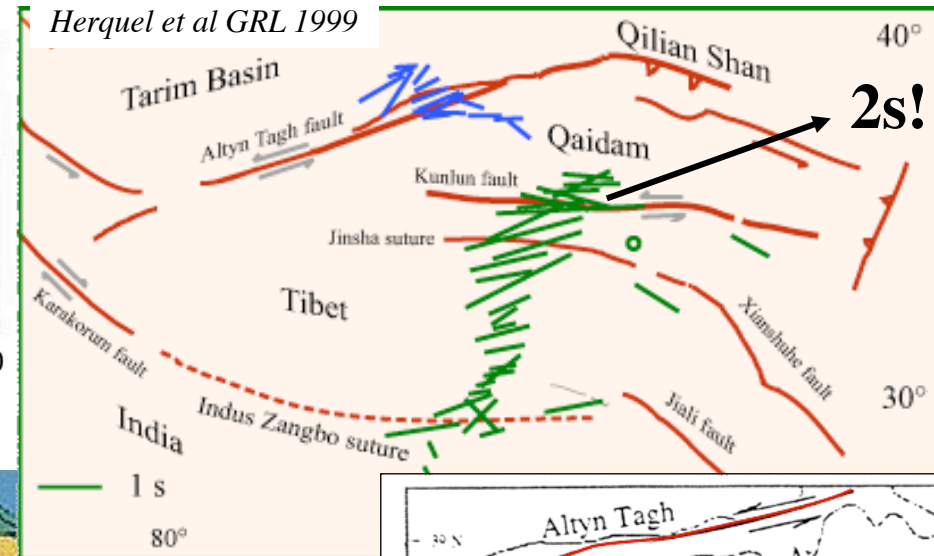
***Shear wave splitting:
fast S wave polarized // to the fault, strong delay times***

Do these large-scale faults penetrate in the mantle?

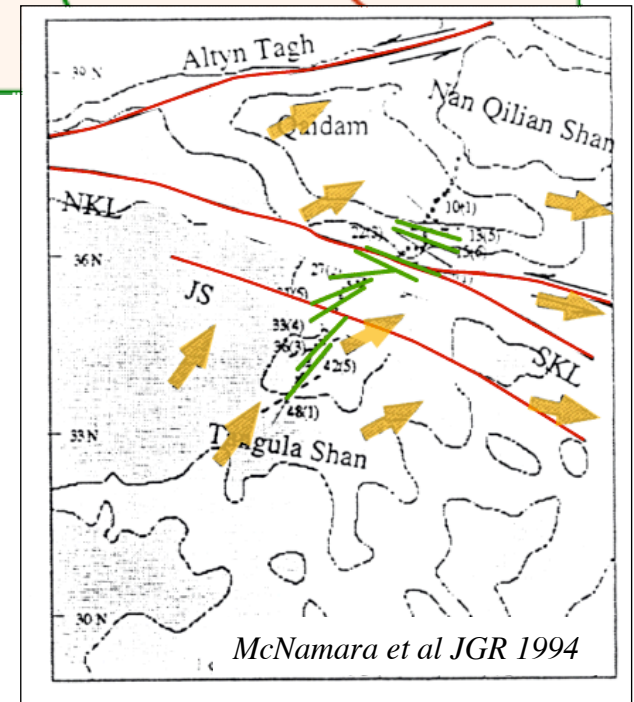
Active faults: Himalaya



Wittlinger et al. Science 1998



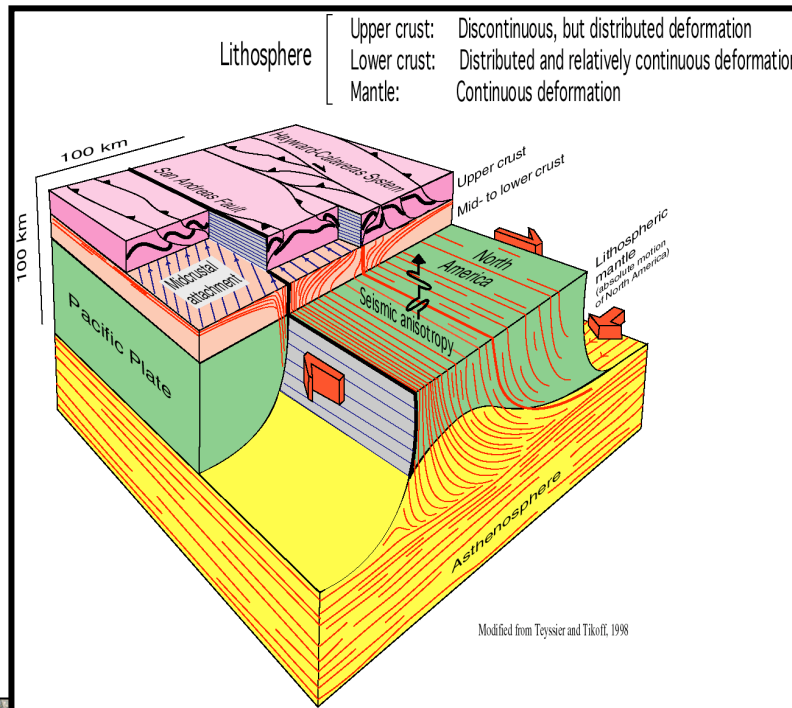
Montpellier



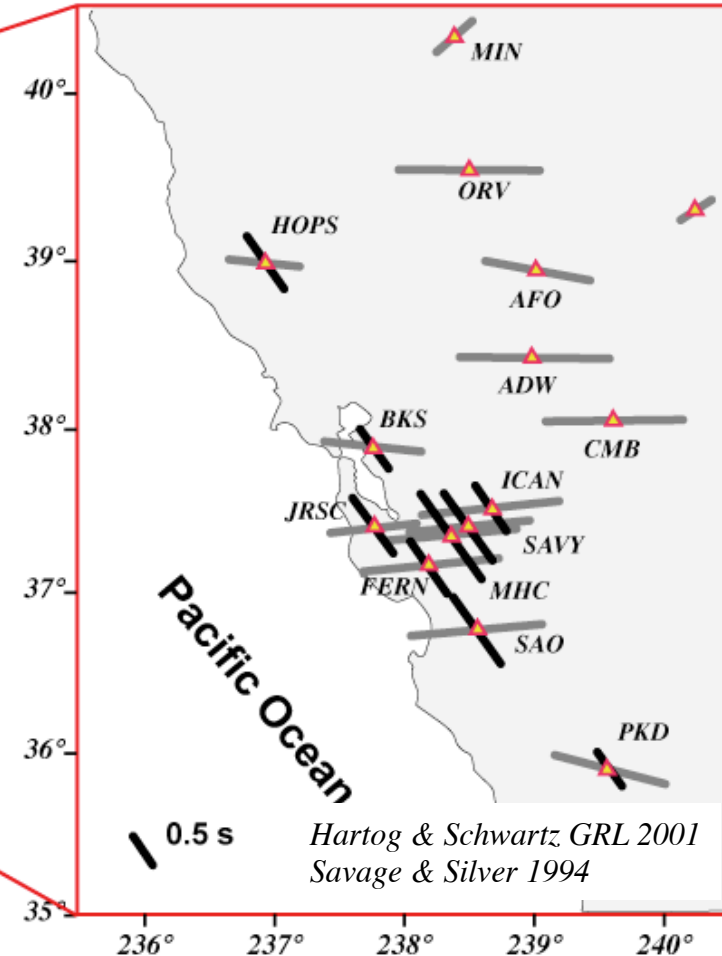
McNamara et al JGR 1994



Do these large-scale faults penetrate in the mantle?



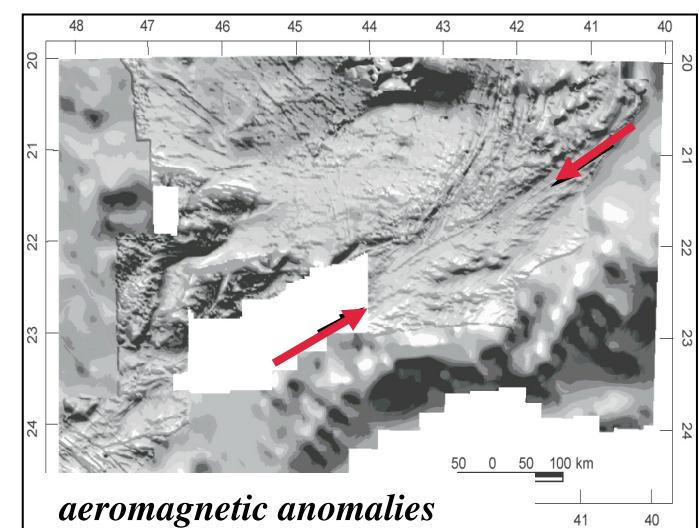
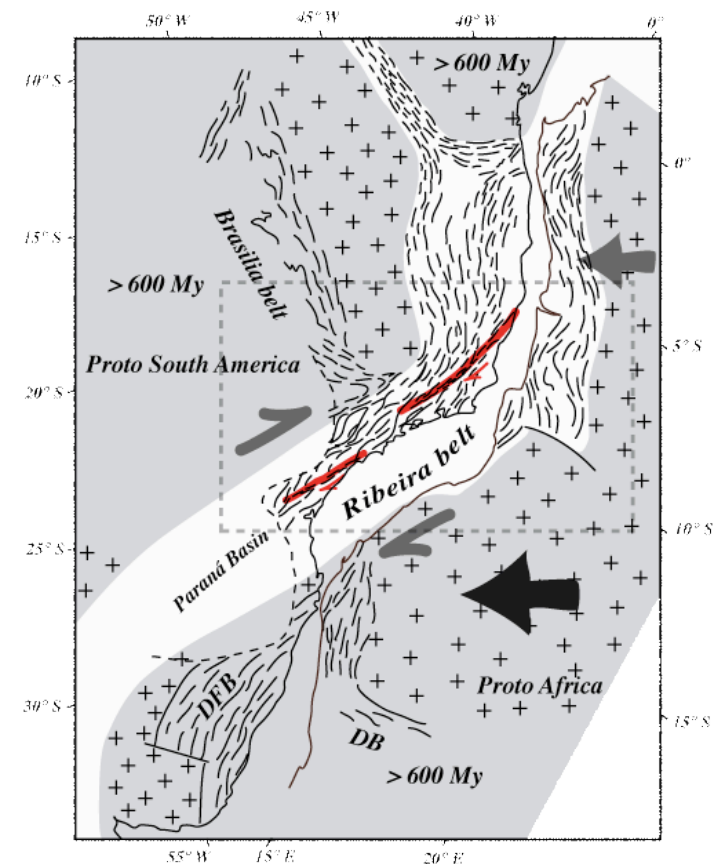
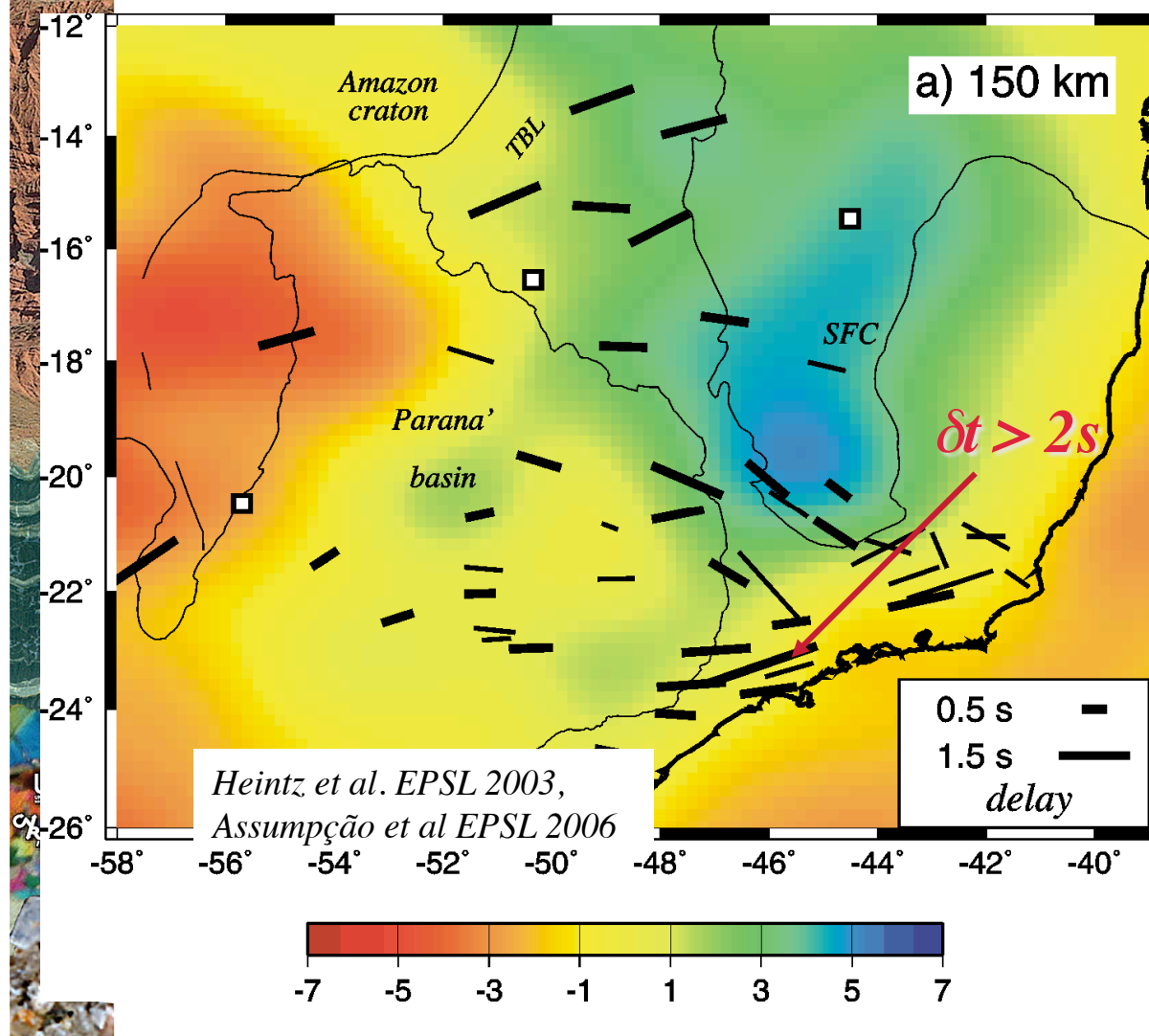
Active faults: San Andreas



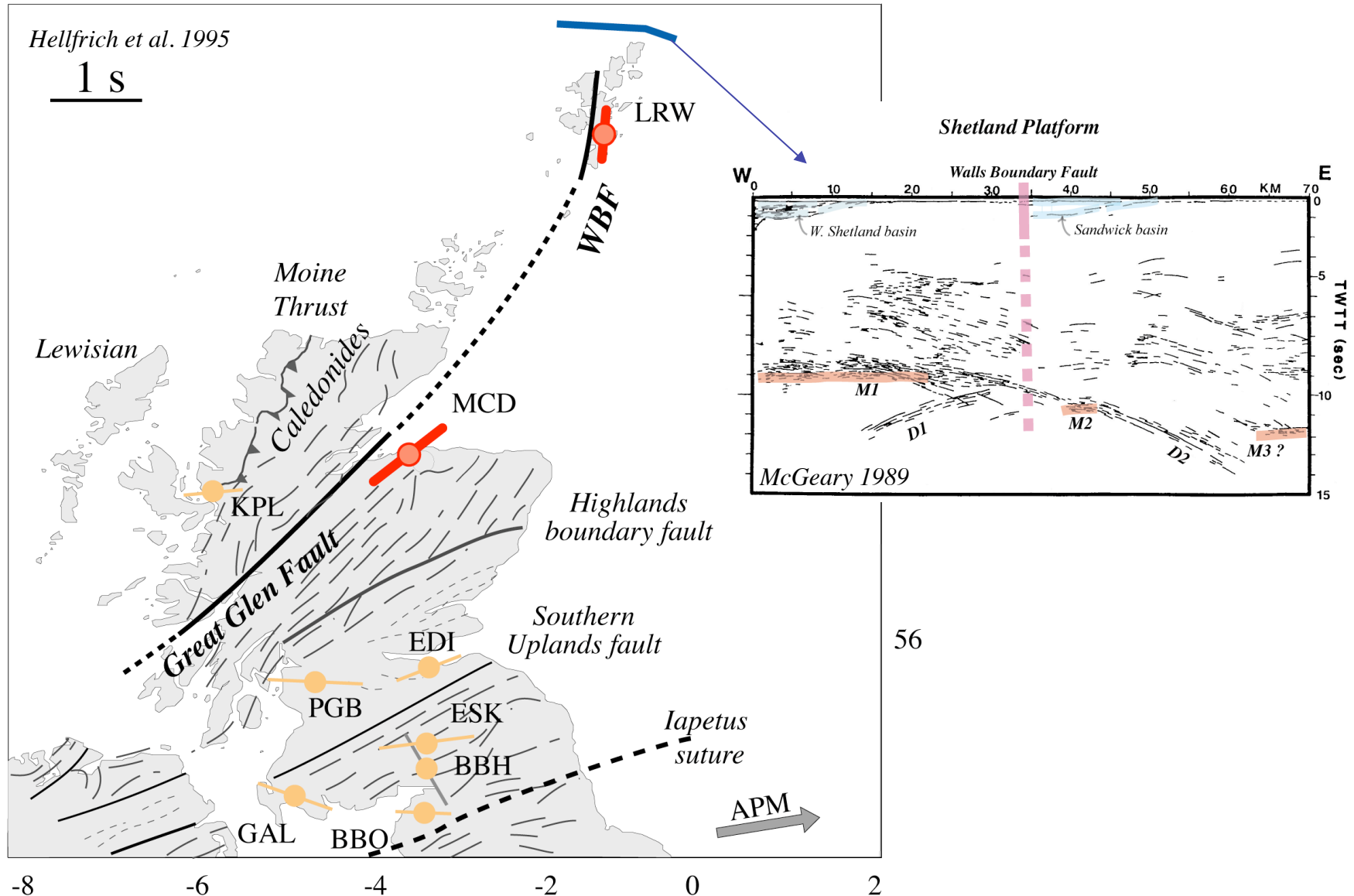
upper anisotropic layer: lithospheric mantle flow // fault



Evidence of coherent crust-mantle deformation: Neoproterozoic wrench faults SE Brazil



Evidence of coherent crust mantle deformation? fossil wrench faults: the Great Glenn fault



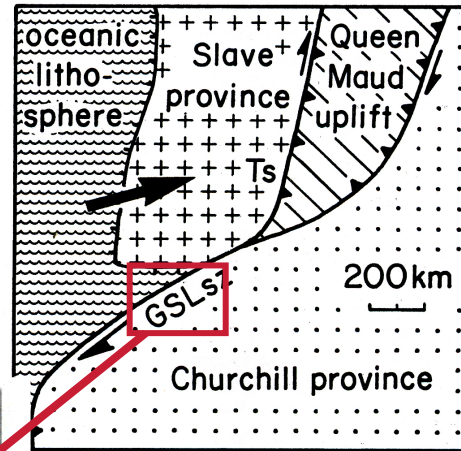
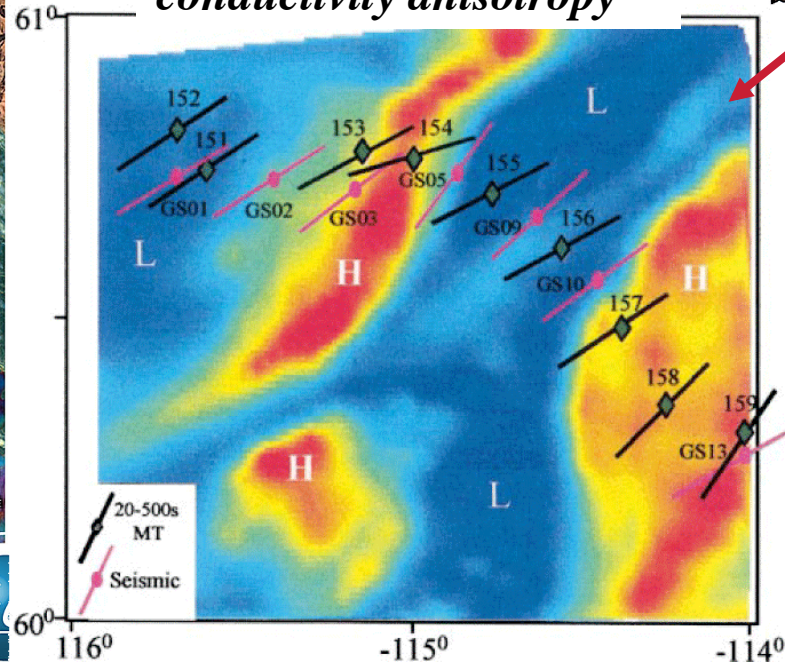
Evidence for coherent deformation of the crust & mantle

Proterozoic Great Slave Lake shear zone

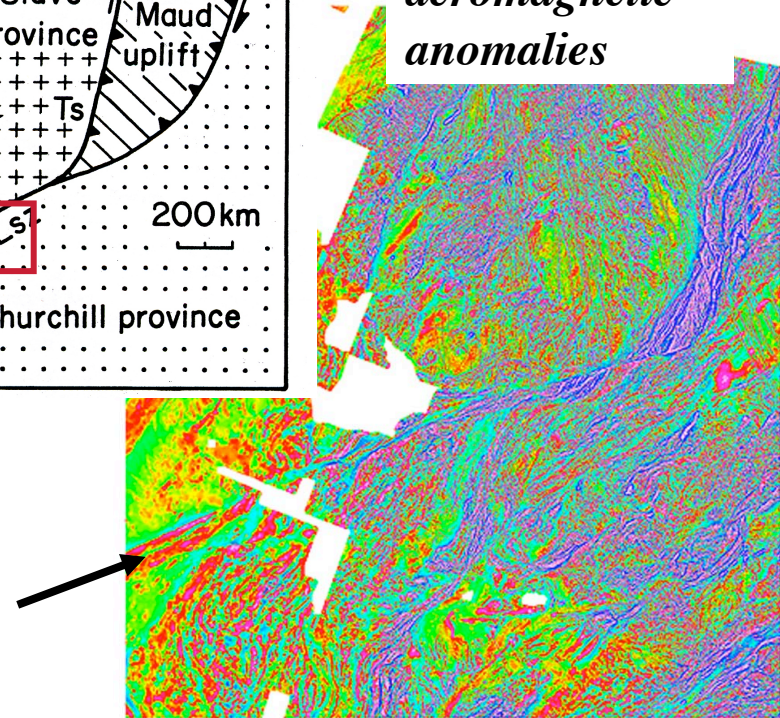
width at the surface >10km
(up to 25km?)

Hanmer et al. 1992; Hoffman 1987

seismic & electrical
conductivity anisotropy



aeromagnetic
anomalies



Canada Grid_ 200m - 1st Vertical Derivative
Commission géologique du Canada

Wu et al. 2002

North Anatolian Fault: no associated anisotropy

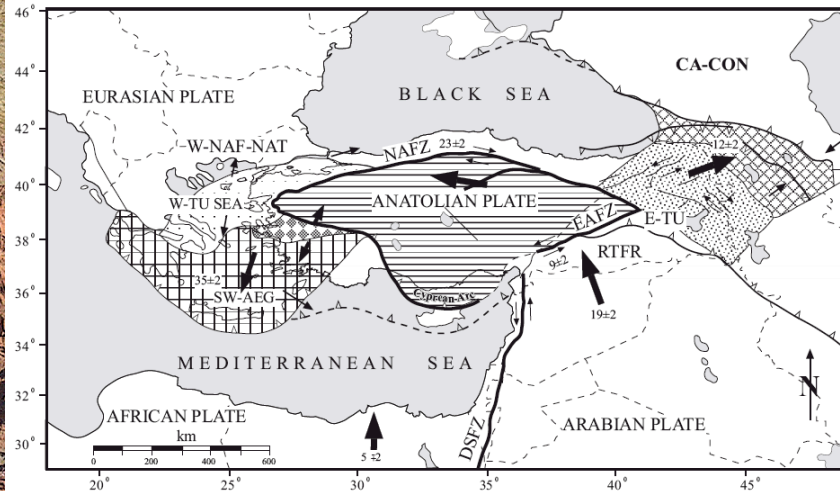
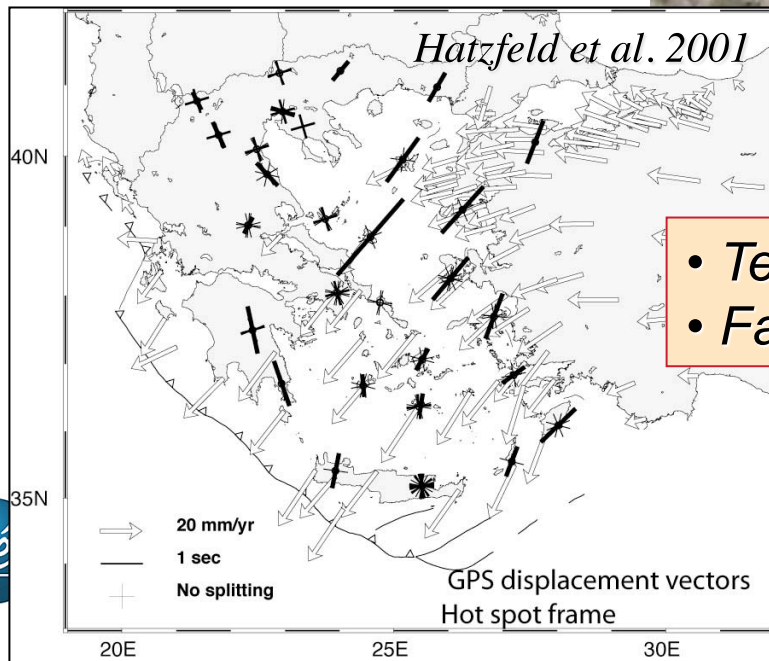
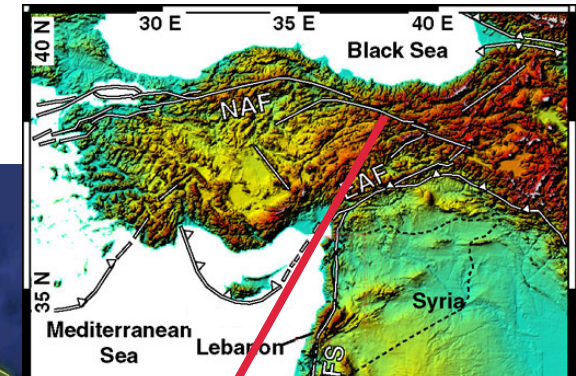
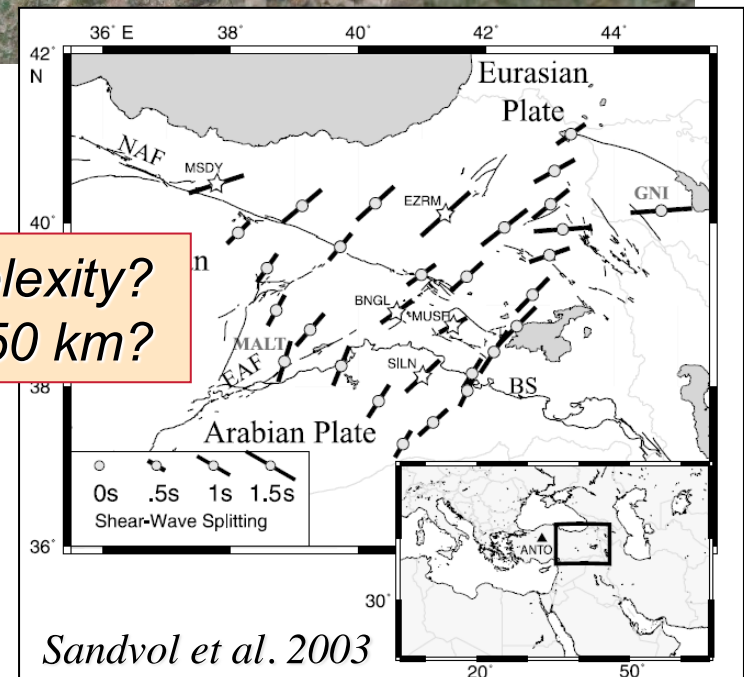


Figure 1. Kinematics of Turkey and surrounding area. Hatched areas show the coherent motion of specific zones of deformation. Heavy arrows indicate the generalised regional motions in mm/yr (after McClusky *et al.* 2000).



- Tectonic complexity?
- Fault width < 50 km?

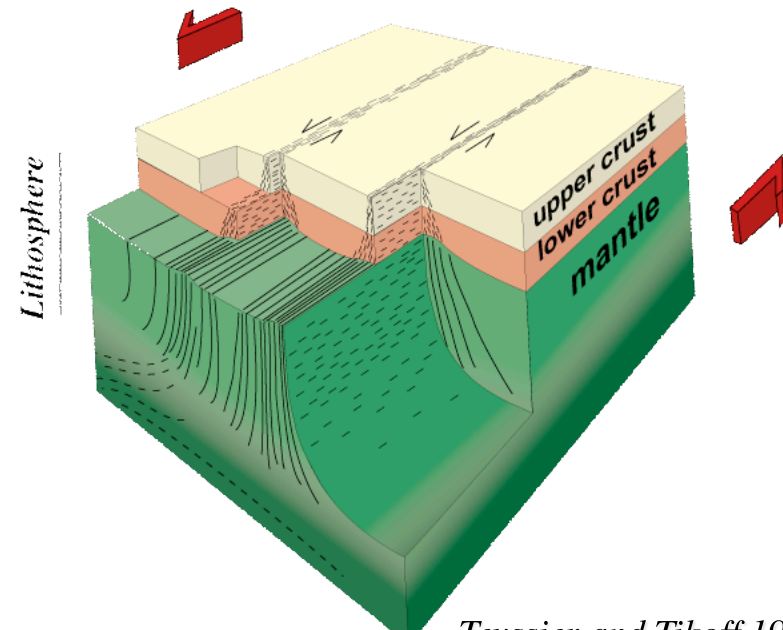


Sandvol et al. 2003

Continental-scale wrench faults in the deep crust & mantle

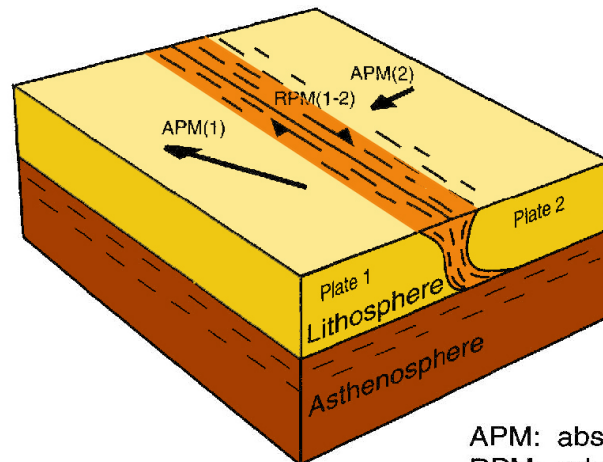
Seismic anisotropy:

- Most (but not all?) large wrench faults deform the lithospheric mantle
- Fault width in the mantle $> 50\text{km}$ ($w < 50\text{km}$: no related seismic anisotropy)
- Delay times $> 2\text{s}$: Coherent flow directions in the lithosphere and asthenosphere

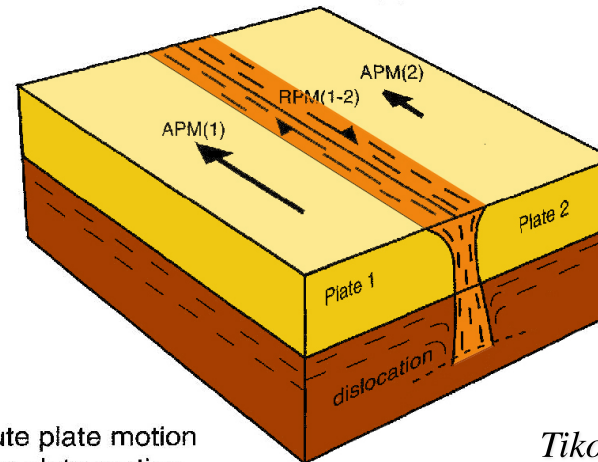


Teyssier and Tikoff 1998
Vauchez and Tommasi 2003

San Andreas type

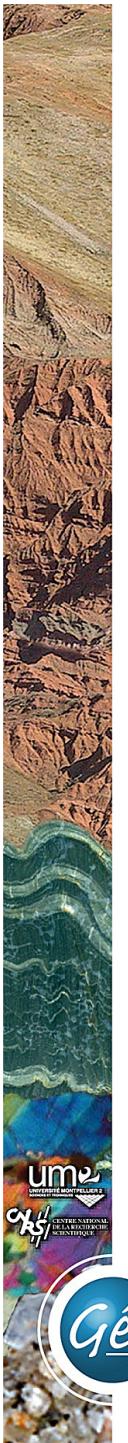


Trinidad type



APM: absolute plate motion
RPM: relative plate motion

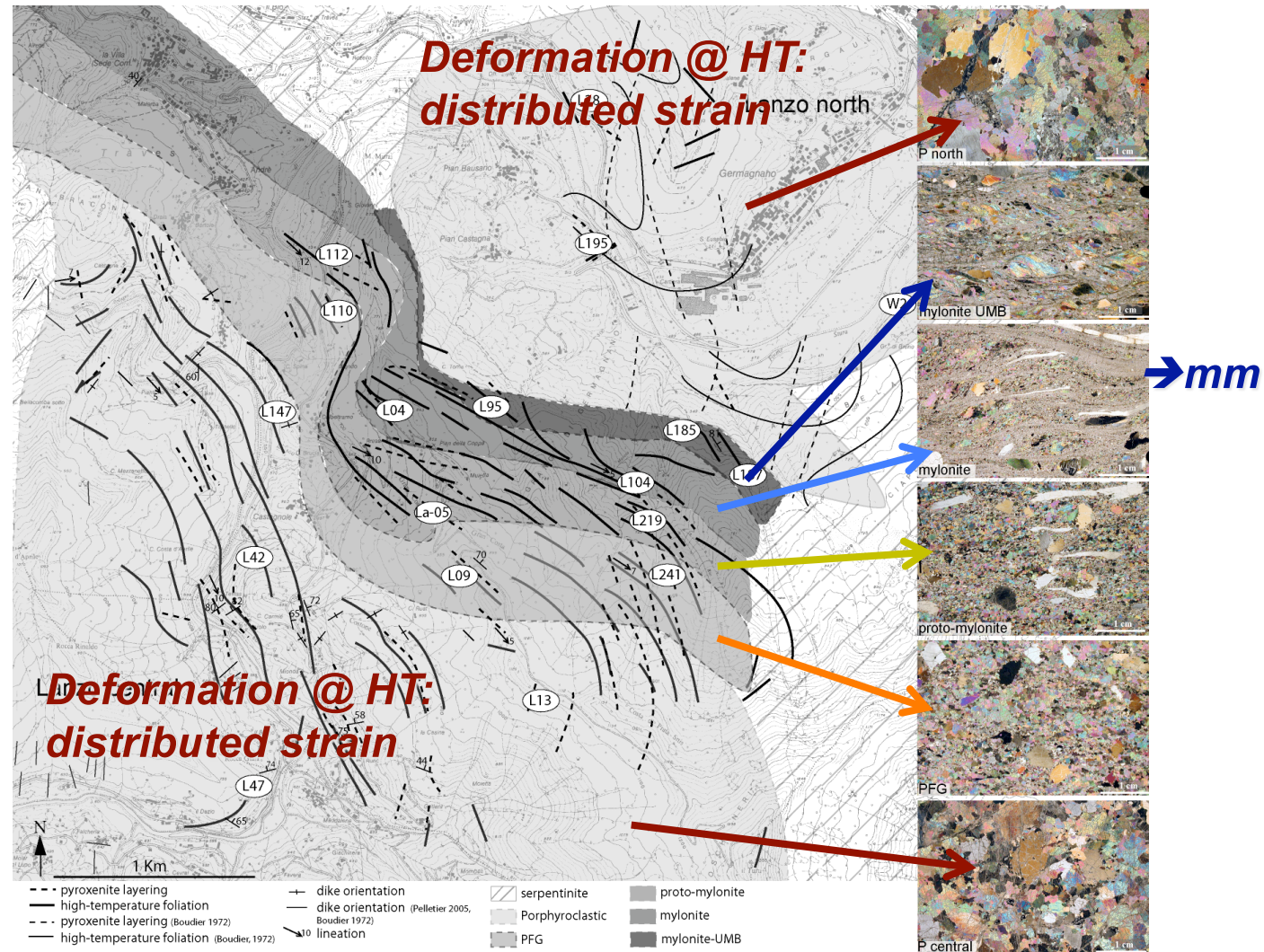
Tikoff et al. 2004



Intrinsic strain localization in the lithospheric mantle?

Lanzo

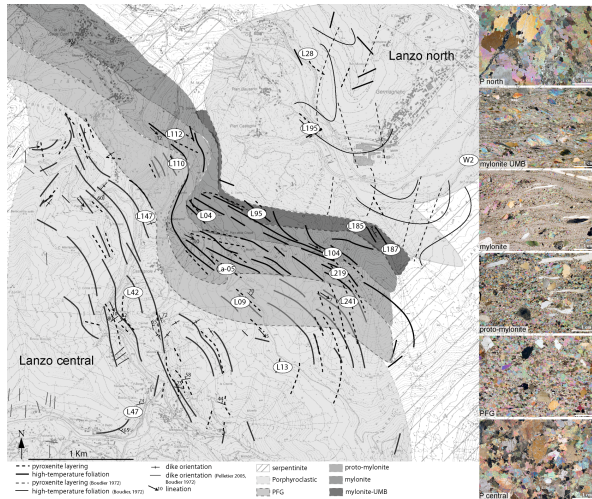
For more details see poster by M.A. Kaczmarek!



deformation under decreasing P,T conditions (>1000-750°C, <1GPa):

- **progressive strain localization** (1km → 600m → 300m → <100m)
- **grain size reduction** (higher deviatoric stresses & strain rates)
- **when & where?**

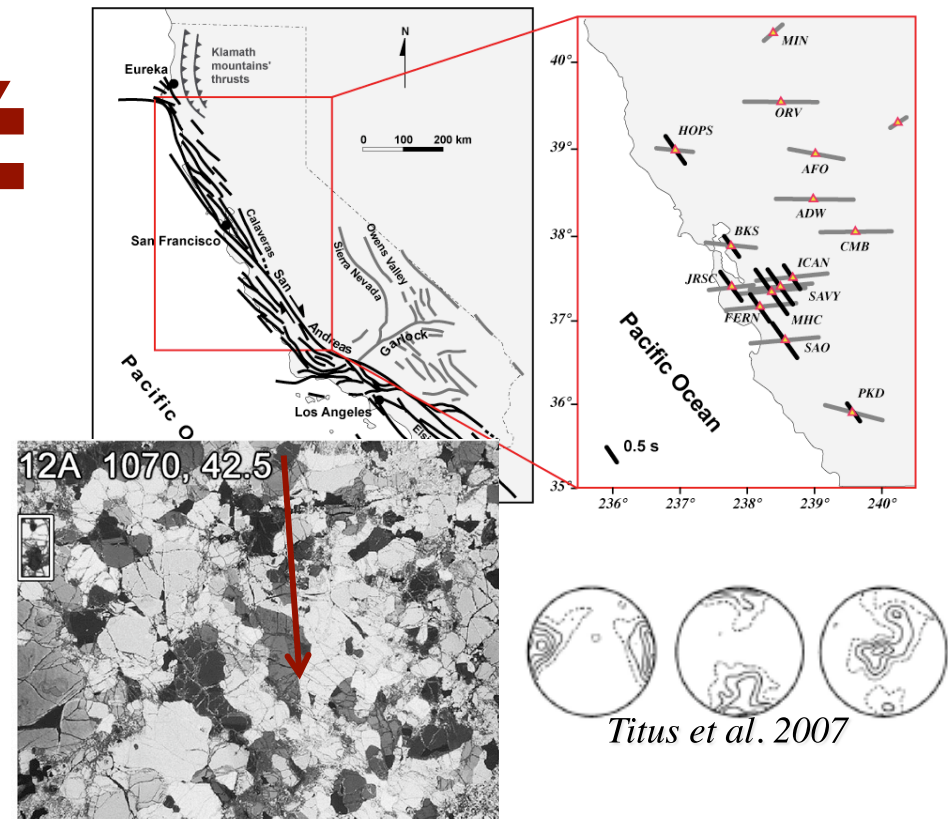
Shear zones in the mantle



≠

- Strong seismic anisotropy in a ≥ 50 km wide zone beneath active and fossil strike-slip faults
- Coarse-grained mantle xenoliths

- Mylonitic zones in the mantle ≤ 1 km
- weak olivine CPO = weak anisotropy



Titus et al. 2007

How does the uppermost lithospheric mantle deform?



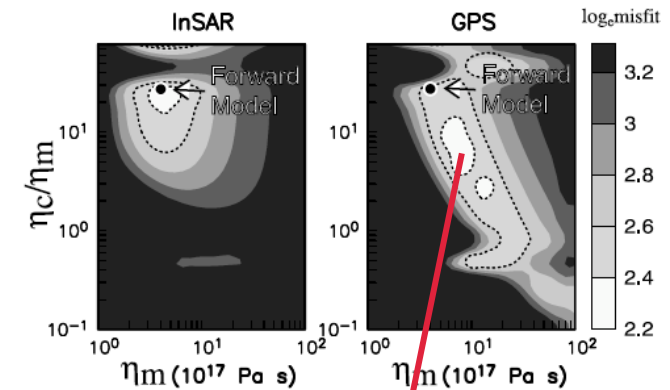
Open question : Rheology of the uppermost mantle

✓ Modeling of post-seismic strain:

→ uppermost mantle viscosity $\ll$ predictions
based on HT lab experiments

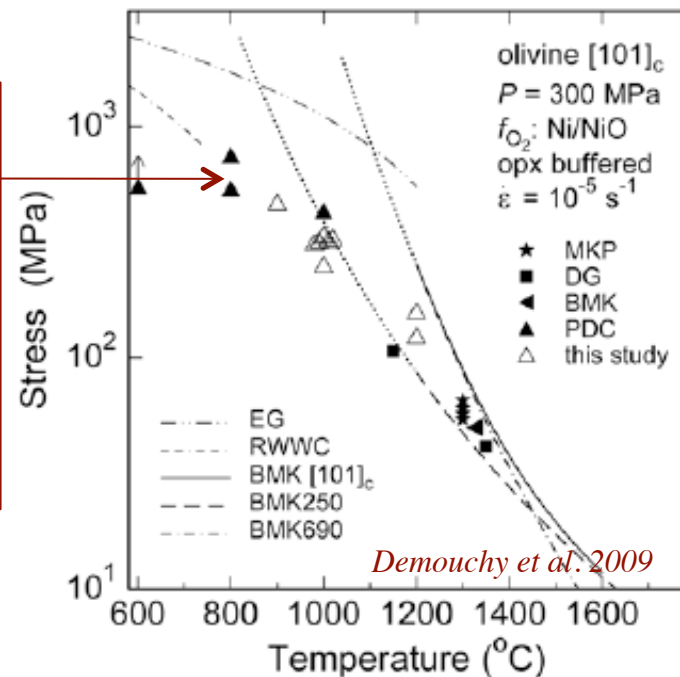
✓ Clear seismic anisotropy signature, while
LT, high stress deformation of peridotites

→ weak CPO & anisotropy!

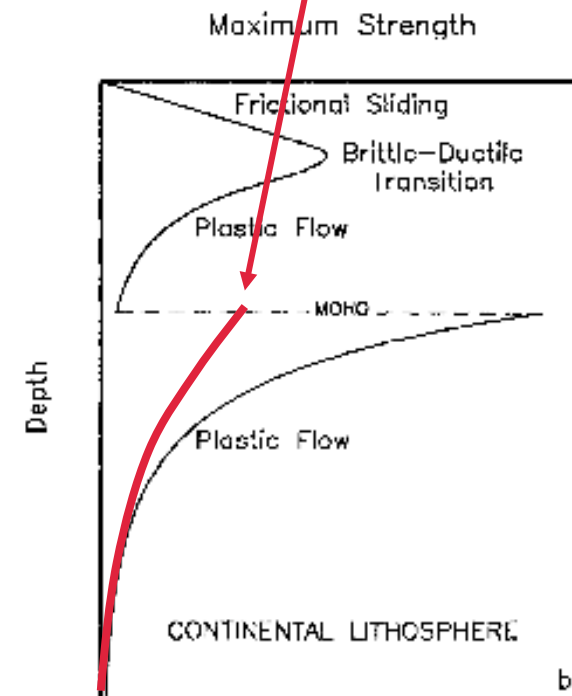


Pollitz et al. 2001

New experiments:
Olivine @ 900-800°C
1-2 orders of
magnitude weaker
than predicted by
extrapolation
of HT power laws!



Demouchy et al. 2009

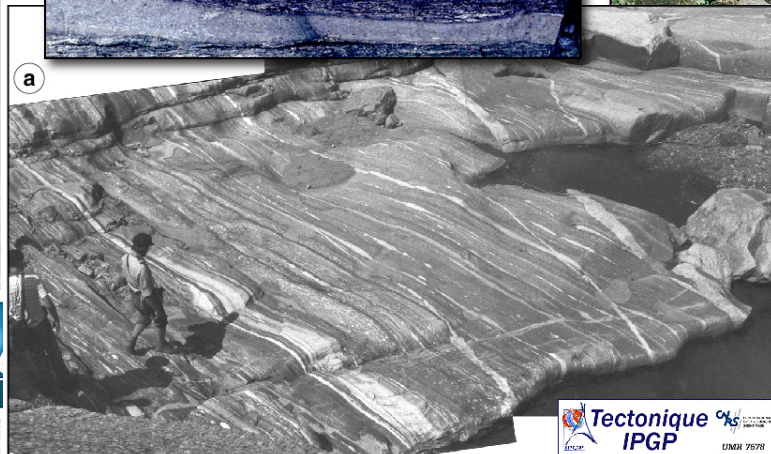
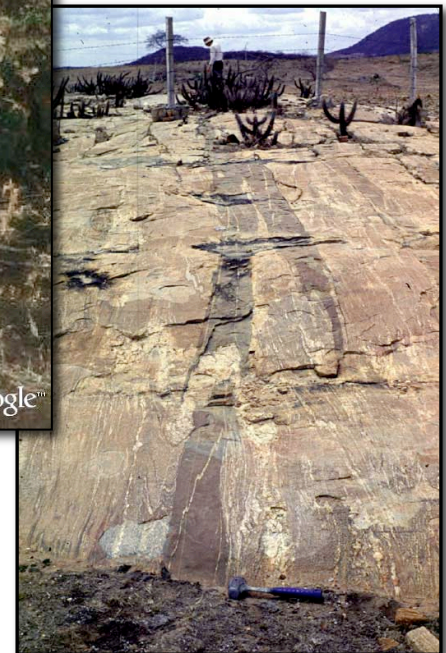


Which processes produce strain localization processes in the ductile deep crust and mantle?

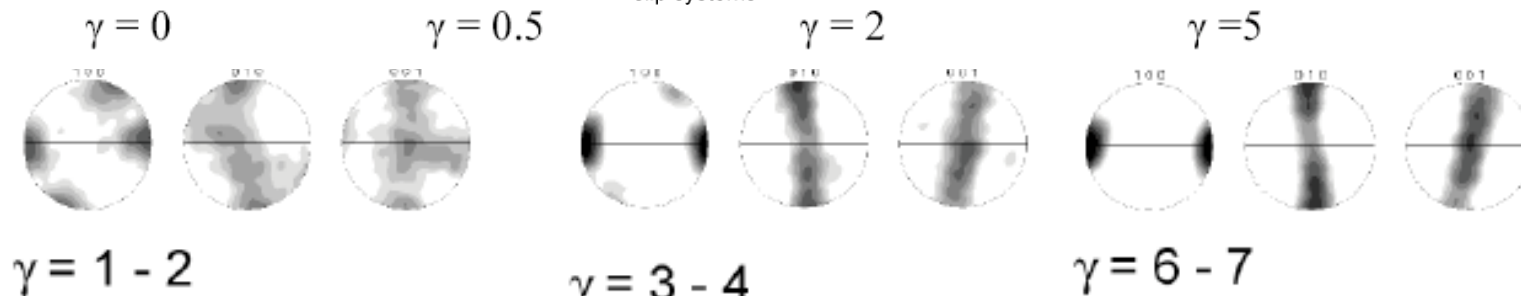
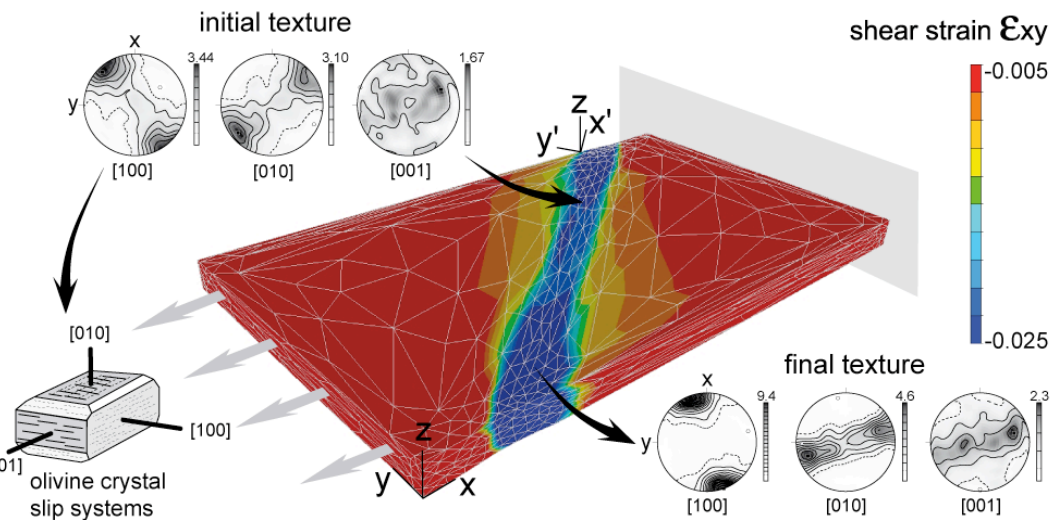
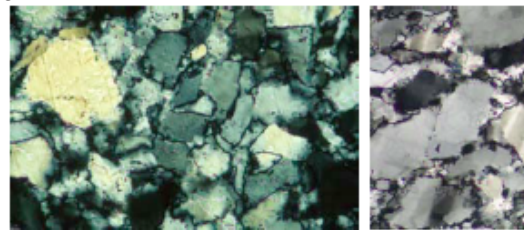
At HT intrinsic strain softening processes not very effective...

How to weaken a ductile fault?

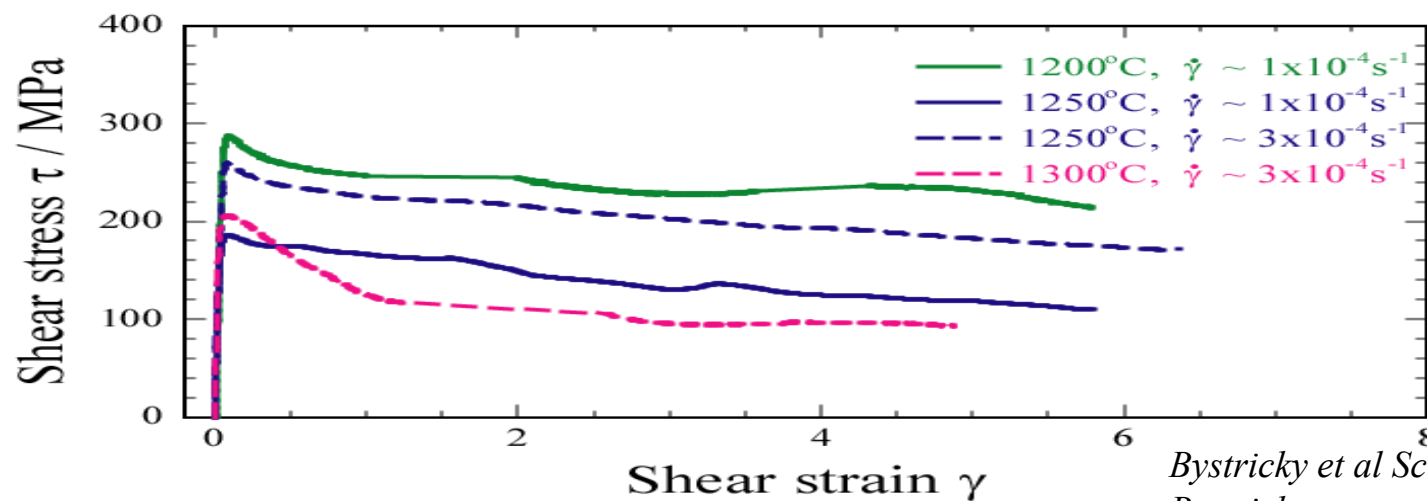
- *Magmas? Other fluids?*



Olivine CPO-induced r Strain wea



Dry Olivine FeFeO

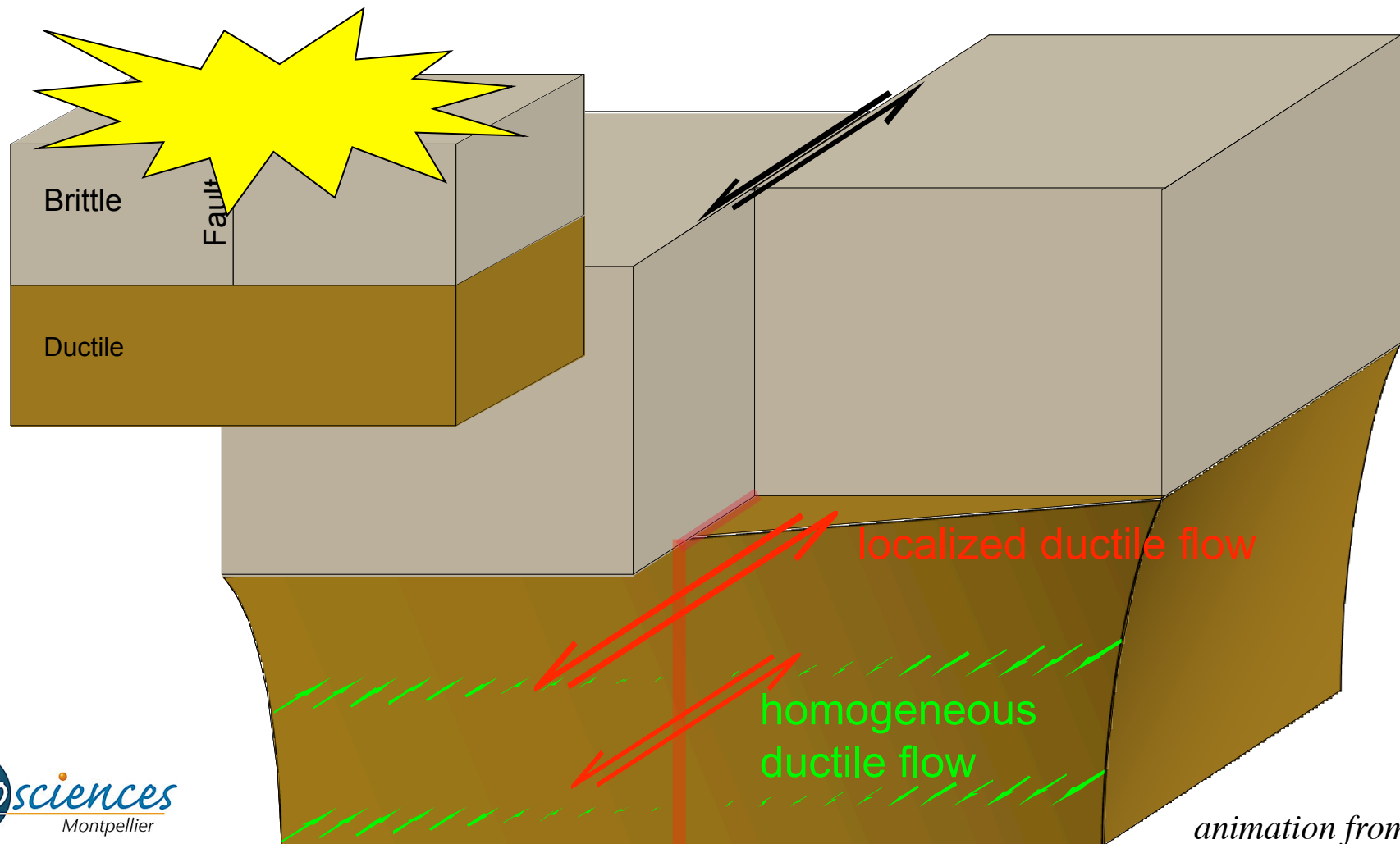


Bystricky et al Science 2003
Bystricky, pers. commun.



Localization controlled by the brittle crust?

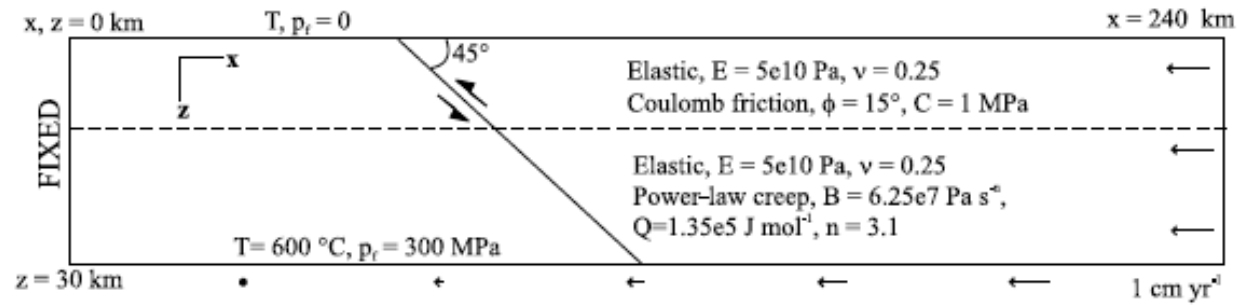
- *Effect of large earthquakes on strain localization in the ductile regime?*



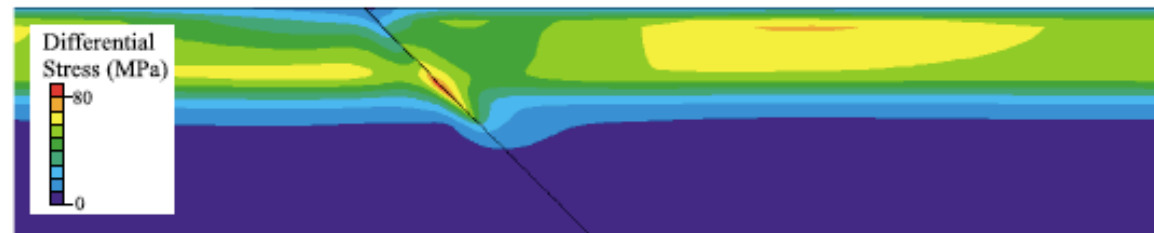
animation from
Laurent Montési



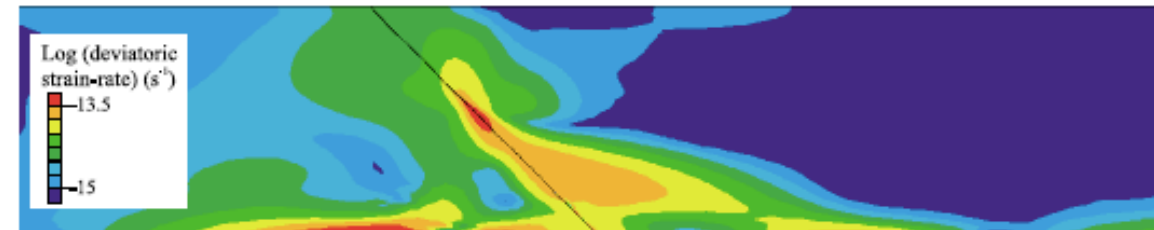
(a) Initial and boundary conditions



(b) Differential stress after steady-state phase



(c) Log (strain-rate invariant) immediately after faulting, 1 seismic cycle



(d) Accumulated creep strain after 8 seismic cycles

